Joint Inversion of Rayleigh Wave Phase Dispersion and Ellipticity and Receiver Functions for crustal S-wave velocity structure of Northern Taiwan

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Introduction



Topography and Bathymetry

Figure 1. Formosa Array brodband stations distribution. Black triangles indicate total 149 seismographs deployed in the northern Taiwan with spacing ~5km.

reduction (Huang et al., 2021).

• The detailed shallow crustal S-wave velocity structure (<10km) phase velocity and ellipticity (Liu et al., 2021).

magma chamber. However, the S-wave velocity properties of the noise correlation data (2-13 s).

- Receiver function. • Horizontal-to-vertical (H/V) amplitude ratio,
- Phase velocity
- from noise cross-correlation function & telesiesmic events. (2-13s)



Figure 6. Example of 1-D MCMC joint inversion result with input data: Receiver function, Phase velocity, and H/V ratio. (a) The inverted 1-D Vs profile in shallow depth (0-10km) (b) Same as (a) but extend to 50 km depth with the full searching space (green dash lines), starting model (red triangles), posterior models (cyan lines), and final 1-D model (white circles) identified. (c) The observed stacking receiver function with uncertainity (black dots and lines) and the receiver function predicted from the starting model (red triangles), posterior models (cyan lines). (d) and (e) Same as (c) but for phase velocity and H/V ratios.

1. The boundaries of velocity contrast (Low-velocity zone & Moho) are detected by receiver functions. 2. The new model revealed more details of mid-to-lower crustal structure after joint inverting Rayleigh wave measurements and receiver functions. 3. The depth range of the magma reservoir (5-15km) matches with Huang et al. (2021) P-wave tomography results (8-20km) but shows a wider geometry (sunny-side-up egg shape). 4. The low-velocity anomalies beneath Ilan plain may imply potential heat sources for geothermal exploration

